Paper VIII



# Fall 2006

# **Outcrops from Every Continent and 20 Countries in 140 Contributions**

Ales of Deep-Water Outerops

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# Formation, Shannon Basin, Western Irelan

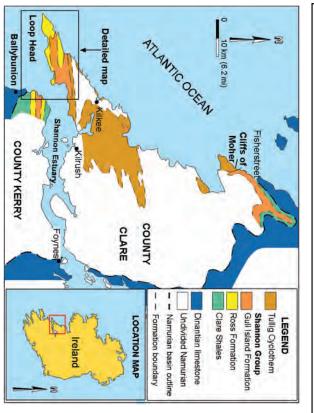
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# Executive Summary

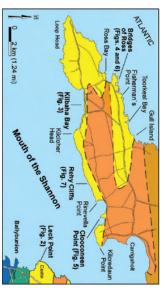
around the world. the stratigraphic succession, make the Ross Formation well suited as an analog for sand-rich turbidite plays in passive margin basins and the basin fill succession begins with the deep-water Clare Shales. These shales are overlain by various turbidite facies of the Ross Clare in western Ireland, and are exposed in sea cliffs along the Shannon Estuary (Figure 1). Carboniferous limestone floors the basin, Formation (460 m [1509 ft] thick). The type of turbidite system, scale of turbidite sandstone bodies, and the overall character of The Upper Carboniferous deep-water rocks of the Shannon Group were deposited in the extensional Shannon Basin of County

and mudstone deposits, and a subsequent vertical aggradation phase with thick-bedded amalgamated deposits. As the channels filled, into channel-spillover belts, and that the belts also followed a sinuous pattern. spilled muds and thin-bedded sands up to 1 km (.6 mi) overbank, followed by thick-bedded amalgamated turbidites that spilled more turbidites spilled farther overbank. Superb outcrops show that thickening-upward packages developed when channels initially individually up to 10 m (33 ft) thick. A few of the upper channels have an initial lateral accretion phase with interbedded sandstone sandier upwards, although there are no small scale thickening- or thinning-upward successions. The upper 290 m (950 ft) consists the same stratigraphic interval but at different locations. This suggests that individual channels and spillover packages were stacked of turbidites, commonly arranged in thickening-upward packages, and amalgamated turbidites that form channel fills that are sinuosity. Stacks of channels and spillover packages 25 – 40 m (82 – 131 ft) thick may show significant palaeocurrent variability at close to the channel margins. The palaeocurrent directions associated with the amalgamated channel fills suggest a low channel The lower 170 m (558 ft) of the Ross Formation contains tabular turbidites with no channels, with an overall tendency to become

upwards packages that formed as spillover lobes. Reservoir elements of the Ross system include frontal-splay tabular turbidites, shallow sandy channel-fill deposits, and thickening-



area of County Clare. Colors mark the different Carboniferous formations, Figure 1. Top is a location map of Ireland, with rectangle showing enlarged Peninsula and Ballybunion area. All areas mentioned in the text are shown. and the Ross Formation is in yellow. Below is the detailed map of Loop Head

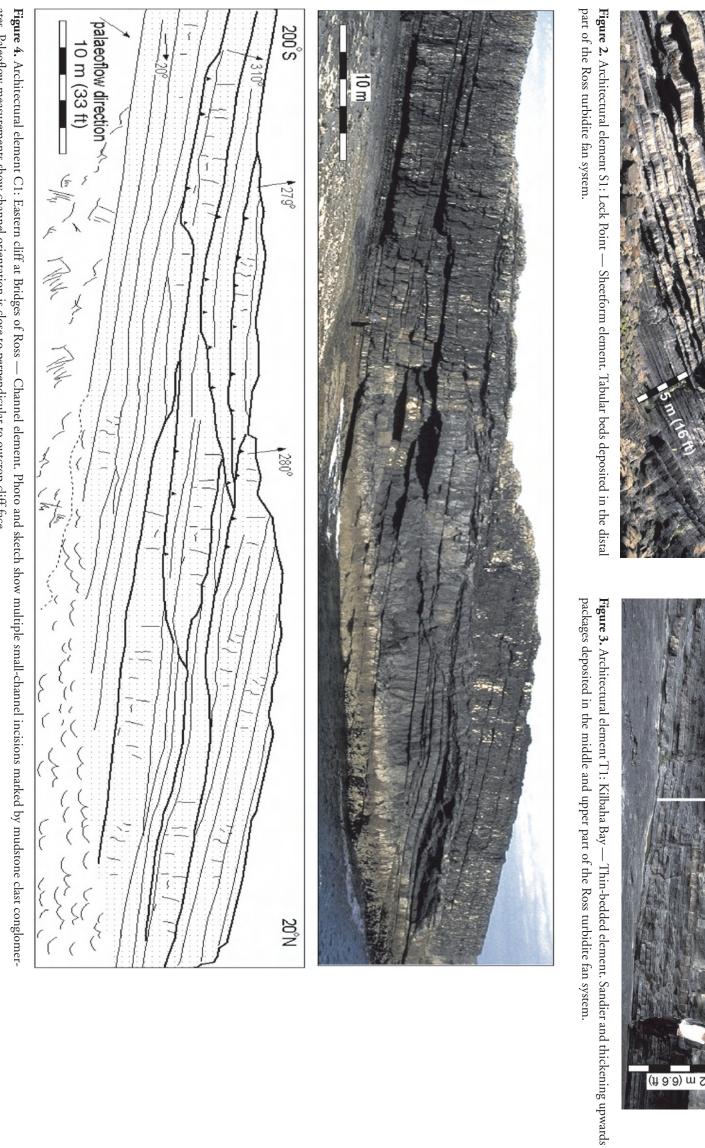


	Outcrop Summary
	Shannon Basin
Location	
Specifics	Shannon Estuary
Region/province/state	County Clare, Western Ireland
Country	Ireland
Formation	Ross Formation
Age	Late Carboniferous (Namurian)
Basin Setting	Extensional basin developed during crustal extension in latest Devonian time. Elongated and confined graben geometry, possible asymmetric controlled by an underlying reactivated of the basis of the basis of the basis of the basis.
	old structural lineament. Deposition on the basin floor along the axis of the basin, onlap- ping the adjacent basin slopes.
Basin Size	ENE-WSW elongation, 20 km <sup>2</sup> (7.8 mi <sup>2</sup> )
General Outcrop Description and Stacking Pat- tern	Sand-rich basin floor fan. Overall sandier upward succession associated with an upward change from tabular sheets to channels and thickening upwards packages interbedded with chaotic units.
Depositional Setting(s) Interpretation Overall Outcrop Dimensions in Panel Length Thickness Average Net/Gross Grain size range Other	Progradational sand-rich fan system with frontal splays in the lower part, to channels with associated overbank lobes interbedded with few large scale slumps and slides in the upper part. The vertical development is gradual, with thin marine bands the only basin-wide cor- relation markers. The frontal splays have a random stacking pattern, the channels are shal- low, sand-filled, have low-sinuousity with some lateral migration and occur as both single and multistory aggradational with a low relief erosional component. Channels occur both isolated and stacked in successions up to 30 m (98 ft) thick and more than 500 m (1640 ft) wide. Overbank lobes are composed of the thin-bedded architectural element and show internal upward sandier and bed thickening trend from mudstone to interbedded thin sand- stone and mudstone beds to amalgamated massive sandstones. Slides and slumps are derived from the lateral basin slopes interfingering with the proximal parts of the fan system. - 23 m (75.5 ft) - 460 m (1509 ft) - 460 m (1509 ft) - Very fine to fine grained sandstone Superbly exposed in sea-cliffs along the Shannon Estuary and the County Clare coast. Con- tinuous unbroken vertical sections more than 400 m (1312 ft) thick can be observed, and

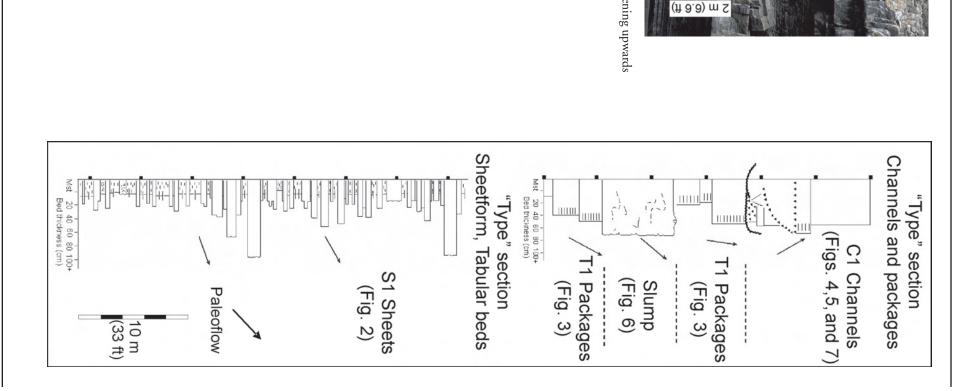
Sheet-form Architectural Elements	>300 m (>985 fr)
Thickness	145 m (475 ft)
Net/Gross	50 - 80%
Outcrop orientation for this element	315 – 135° (Figure 2)
Average paleocurrent	75° (Figure 2)
Typical facies succession	Bouma A
Sand/Conglomerate Bed Architecture	
Bed length range	10->300 m (32->985 ft)
Bed length average	>300 m (>985 ft)
Bed thickness range	0.1–1.27 m (0.3–4.2 ft)
Bed thickness average	0.14 m (0.46 ft)
Texture - grain size range, average, and sorting	Very fine to fine, well sorted
Shale Bed Architecture	
Bed length range	10->300 m (32->985 ft)
Bed length average	>300 m (>985 ft)
Bed thickness range	0.01–5 m (0.03–16.4 ft)
Bed thickness average	~0.10 m ~(0.32 ft)
Channel-form Architectural Elements	
Channel form	Partial
Width	100–200 m (328–656 ft)
Thickness	5–10 m (16–32 ft)
Aspect ratio: Width/thickness	20
Net/Gross	9920
Outcrop orientation for this element	200–020° (Figure 4)
Average paleocurrent	300° (Figure 4)
Typical facies succession	Massive sand, mud-conglomerate at base
Channel infill bedding architecture	Massive amalgamated channel fill. Channel margin onlap. Occasional lateral accretion.
Sand/Conglomerate Bed Architecture	
Bed length range	NA
Bed length average	NA
Bed thickness range	~0.1–0.8 m (~32–262 ft)
Bed thickness average	~0.25 m (~82 ft)
Texture - grain size range, average, and sorting	Very fine to fine, well sorted
Channel-base shale drape coverage	-75%
Channel-base shale drape thickness	01m (03.3 ft)
Shale Bed Architecture	
Bed length range	NA
Bed length average	NA
Bed thickness range	NA
Bed thickness average	NA

Thin-bed Architectural Element	T1
Length	>250 m (>820 ft)
Thickness	3 m (9.8 ft)
Net/Gross	60%
Outcrop orientation for this element	260–080°
Average paleocurrent	50°
Typical facies succession	Bouma A
Sand Bed Architecture	
Bed length range	10->250 m (32->820 ft)
Bed length average	>250 m (>820 ft)
Bed thickness range	1–100 cm (0.03–3.2 ft)
Bed thickness average	10 cm (0.32 ft)
Texture - grain size range, average, and sorting	Very fine to fine, well sorted
Shale Bed Architecture	
Bed length range	1->250 m (3.2->820 ft)
Bed length average	>250 m (>820 ft)
Bed thickness range	1–50 cm (0.03–1.64 ft)
Bed thickness average	5 cm (0.16 ft)
Analog Reservoirs or Fields	NA





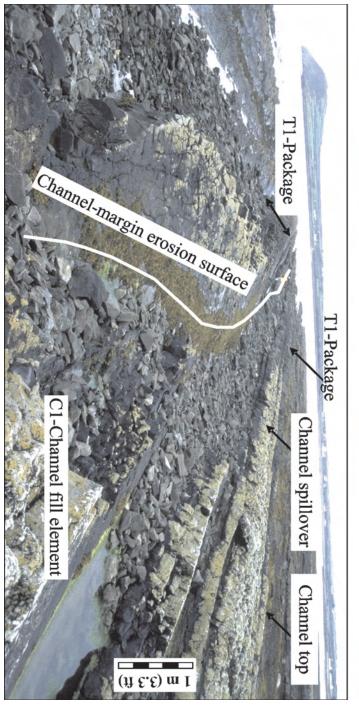
ates. Paleoflow measurements show channel orientation is close to perpendicular to outcrop cliff face.



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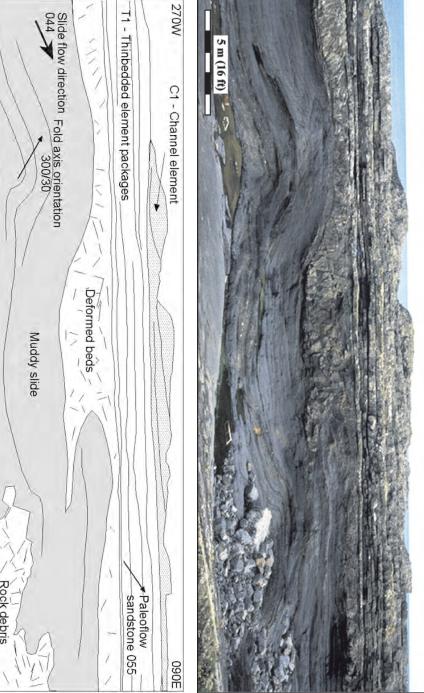
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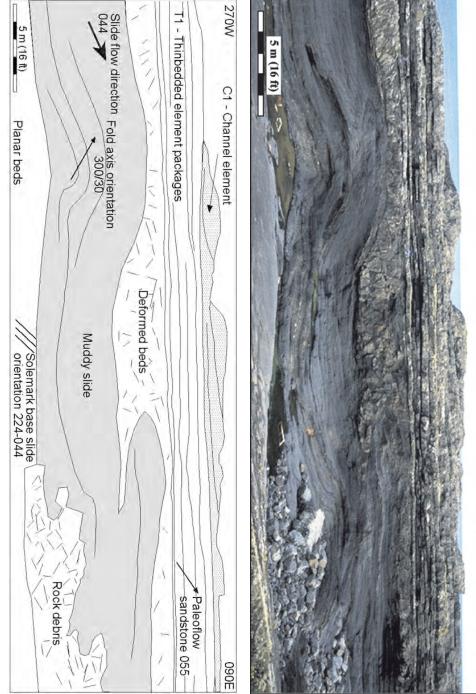


top of the channel fill passes into the top of a thickening upward package towards the west. Channel architectural element to T1-Thin-bedded architectural element. White line shows channel margin surface which cuts down towards the east. Note onlap of channel fill against this erosion surface, and spillover and onlap of higher beds towards the west. The Figure 5. Clooconeen Point — Channel margin and spillover with view approximately to west. Lateral spillover to the west from C1-



about 6-8 m (20 - 26 ft), and the length of the set is about 200 m (656 ft). Dipping beds are interpreted as lateral accretion deposits. recessive intervals are mud-clast conglomerate, these intervals and the sandstones are gently dipping to the left at cliff top. The sandstones pinch out toward the top, and toward the lower downlap surface, giving sigmoidal bed shapes. The thickness of set of dipping beds is Figure 7 (lower left photo). Cliffs below Rehy Hill — Lateral accretion of a C1 - Channel. Regional bedding is horizontal. The dark





Slide are planar thin-bedded sandstones. of photo below, note relatively similar paleoflow direction of the muddy slide and the thin-bedded sandstones above. Below the Ross Thin-bedded element (thickening and sandier upward packages) to vertical and lateral stacked channels (C1: Channel element). Sketch Figure 6. Bridges of Ross — Stratigraphic upward change from the 6 m

(20 ft) thick muddy slide (the Ross Slide) to undeformed T1-